

Groundwater modelling in semi-arid Africa: challenges and lessons

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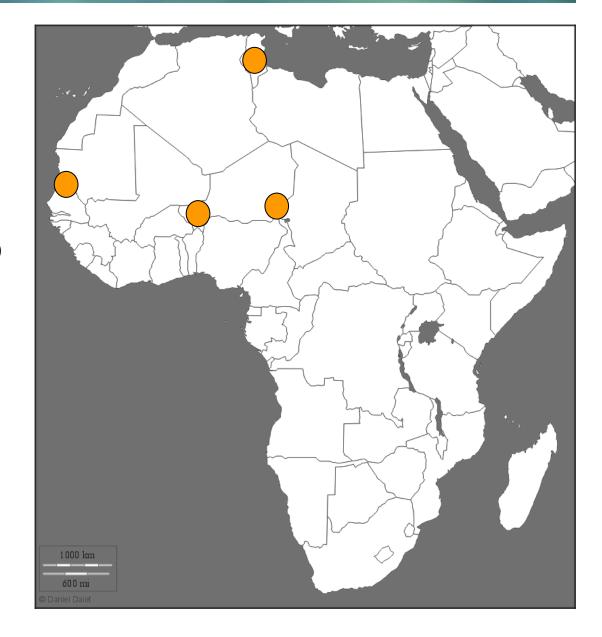


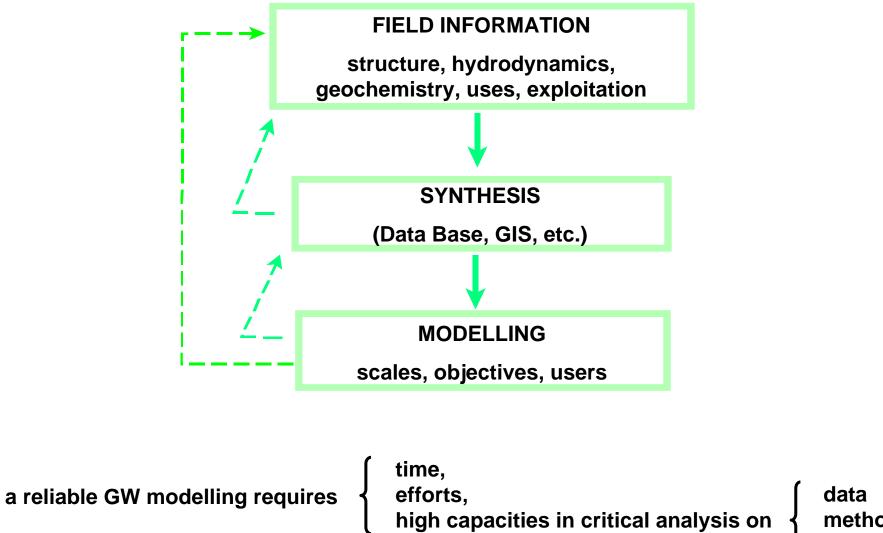
Transboundary Aquifers Cooperation in Africa, Tunis, September 2017

Context

ideas and proposals based on various experiences from hydrogeological studies in semi-arid Africa

Merguellil catchment (central Tunisia) Senegal-Mauritania Basin Iullemeden Basin (SW Niger) Lake Chad Basin and others





data methods results

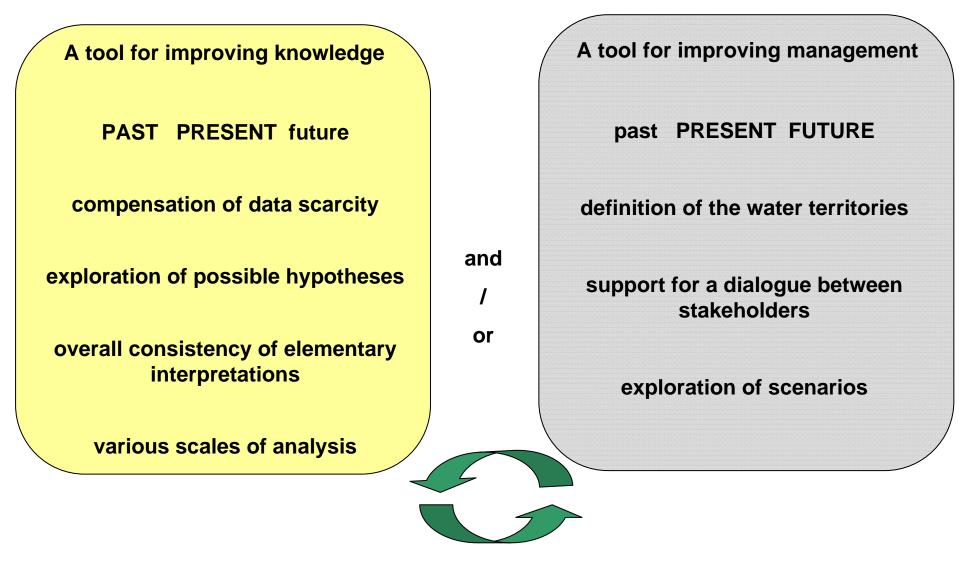
an intellectual construction

integrating a reasonable part of the field complexity aiming to reproduce the behaviour of the socio-physical system according to a given level of knowledge in response to specific questions without a unique solution

an overall quality depending on data and assumptions intrinsic capacities of the model and modeller

> As far as the laws of mathematics refer to reality, they are not certain; and as far as they are certain, they do not refer to reality

> > **A. Einstein** English translation of the conference "Geometrie und Erfahrung", Berlin, 1921



necessary interactions

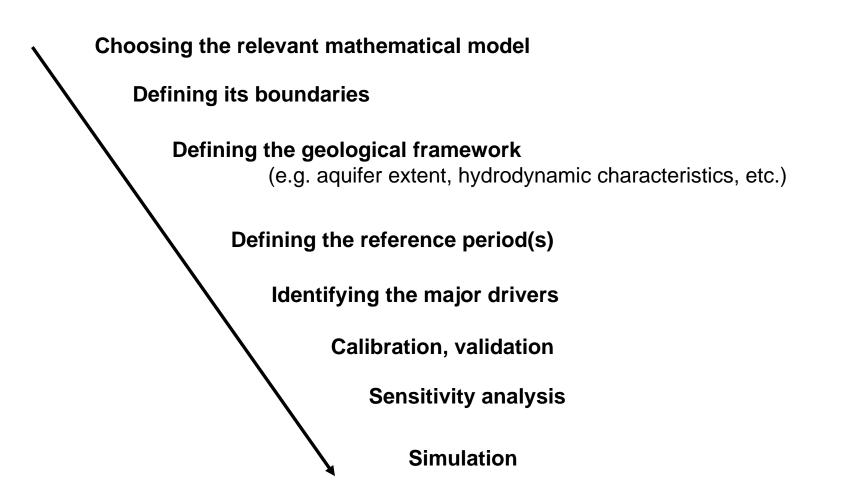
Severe limitation by

quantity, quality, representativeness of data density of information vs. complexity of physical and social phenomena rapidity of temporal changes

Importance of

long-term processes exceptional events rapid changes

Interests of exploring processes that cannot be measured directly in the field



Choosing the relevant GW model

Evapotranspiration Paramètres fixés pendant la procédure de calage Phuie EAU CRT FN R QRMAX DCRT CQR QIMAX Ruissellement RNAP CQI Infiltration

black box and other empirical models



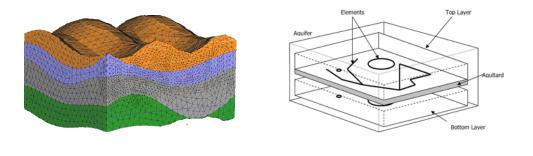
operationality

black box

mathematics adapted to processes

• porous medium

management of data, I/O satisfying level of accuracy skilfullness of the modeller/user



physically based (e.g. Darcy)

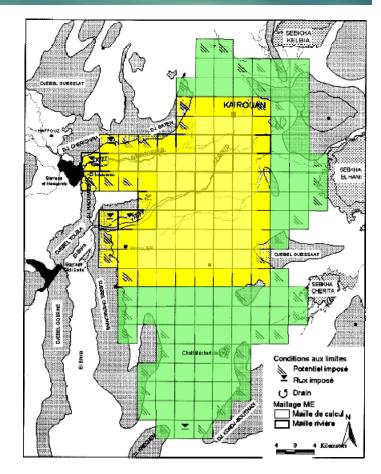
relevant scales in space and over time (data, objectives of knowledge vs. management)

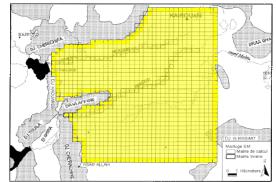
major processes (in the present and future state) requirements vs. neglegts

internal vs. external interactions

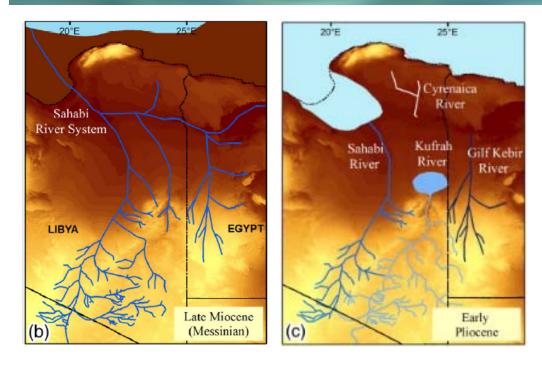
influences of boundaries on calculation

i.e. exhaustiveness vs. relevancy and efficiency



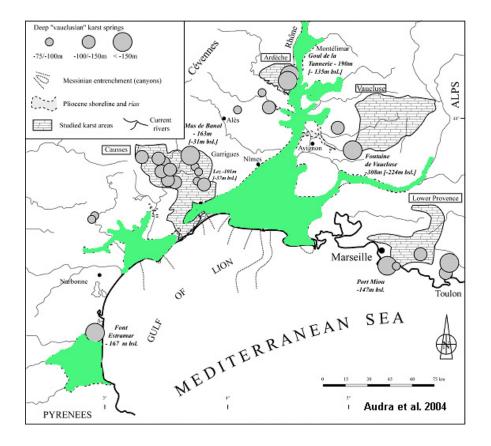


The geological framework

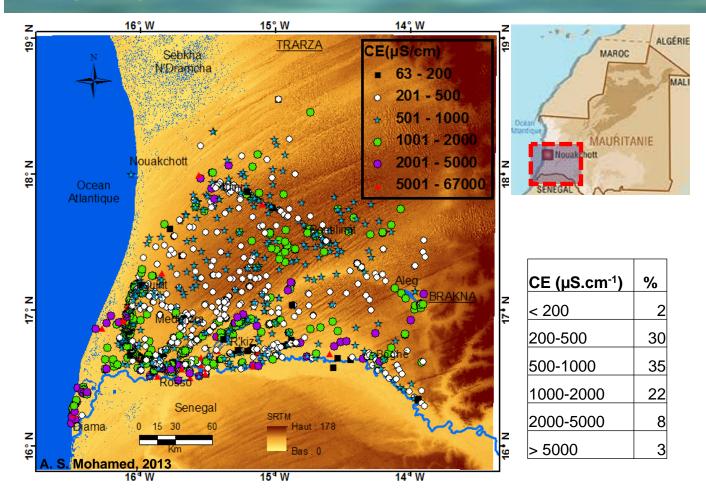


Messinian rias in SE France and karstic springs

The Kufrah paleoriver (Libya)

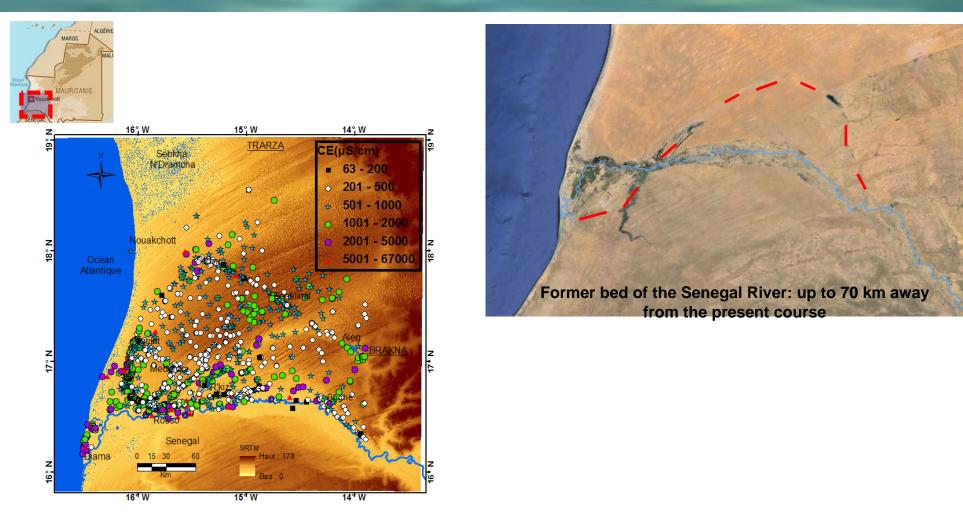


The geological framework (2)



The present distribution of GW salinity is not explained by mineralogical variations in aquifer sediments.

The geological framework (3)

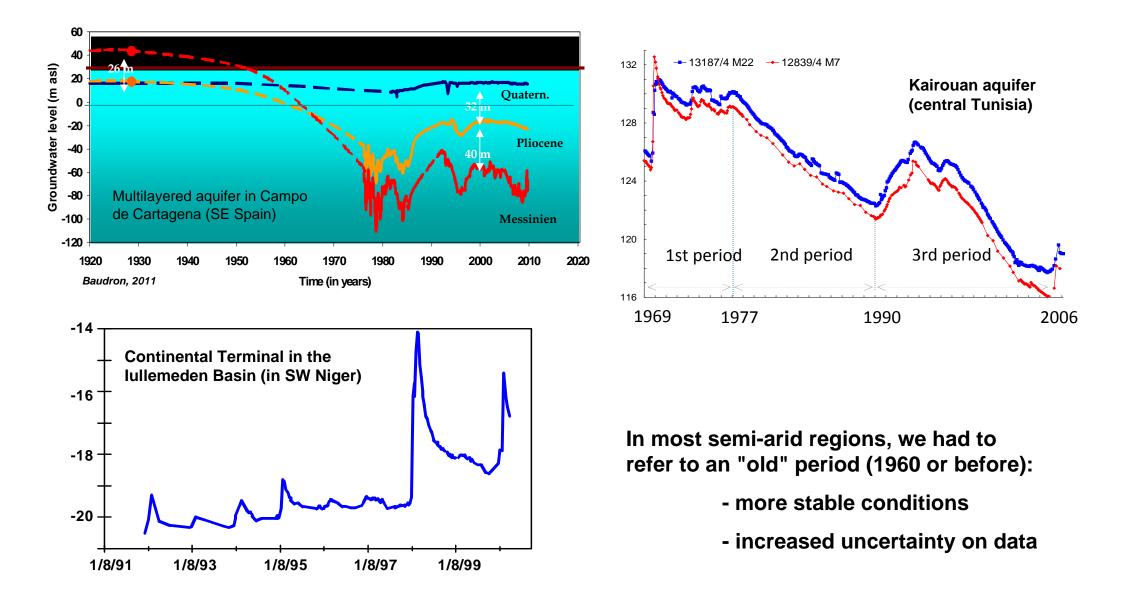


The present distribution of GW salinity is primarily explained by

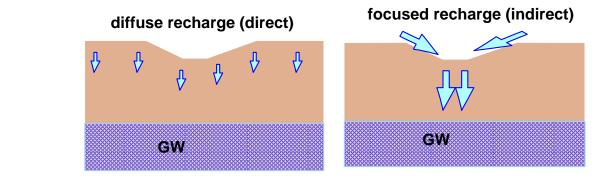
- Holocene eustatic fluctuations, ans seawater intrusion up to 200 km upstream in the Senegal valley,

- changes in the location of the Senegal River bed.

Defining the reference period

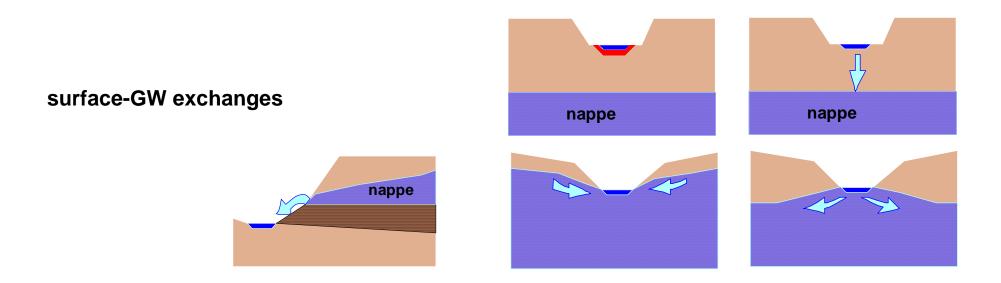


Example of the recharge identification: processes

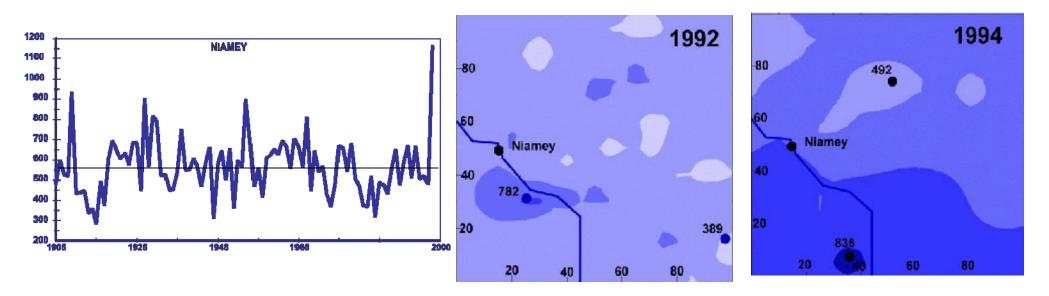


rainfall infiltration

proportion of focused recharge is assumed to increase with aridity

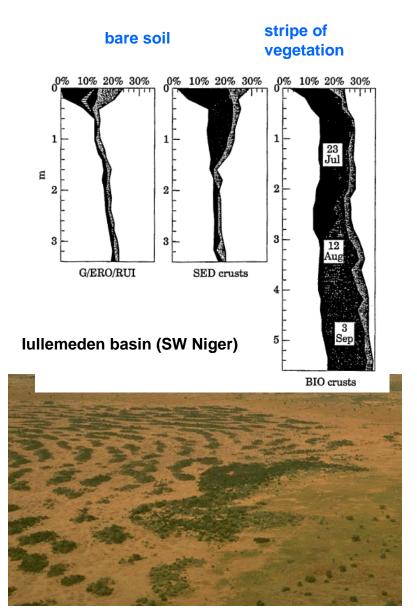


Variations of rainfall in space over time

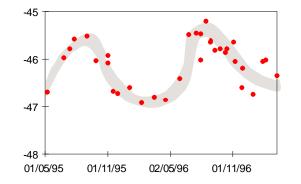


largely exceeding the density of observation networks

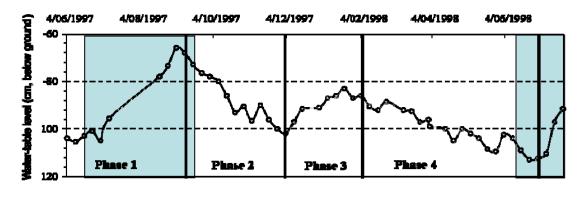
Example of the recharge identification: rainfall



Transformation of rainfall into recharge



seasonal variation

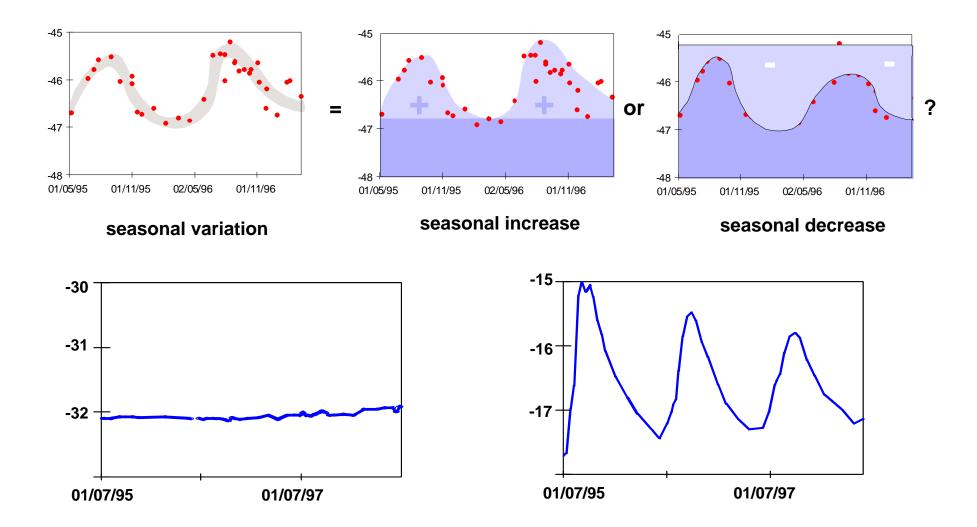


one rainy season, two GW rises per year

(Lake Chad Basin)

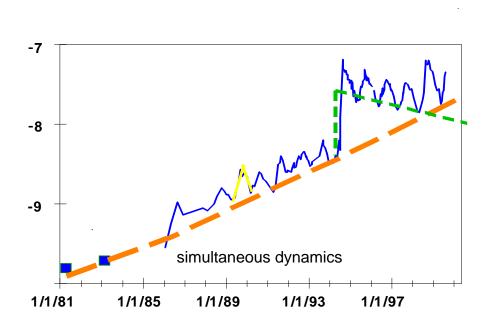
Observations from the lullemeden Basin

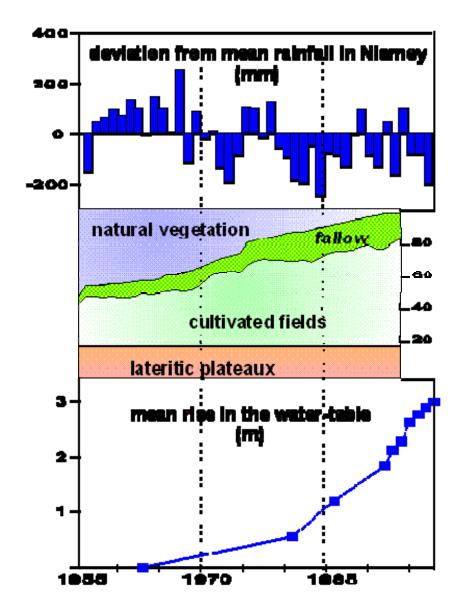
over a very short distance, in a phreatic aquifer open to rainfall infiltration



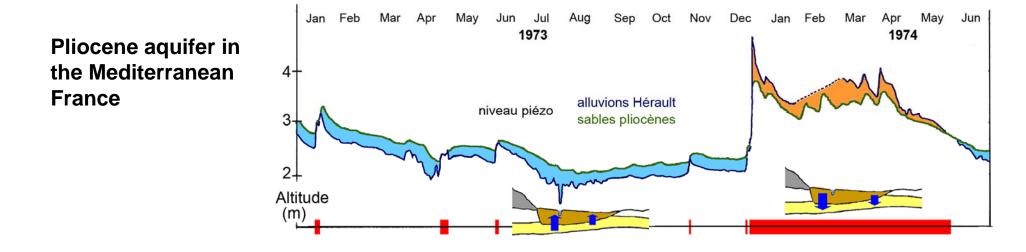
Example of the recharge identification: rainfall

Long term changes in the lullemeden Basin





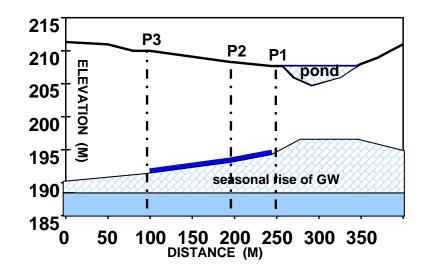
Example of the recharge identification: surface-GW exchanges

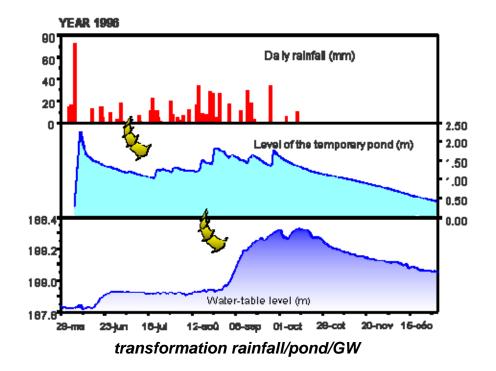


Exceptional event (once in 20 years), with a long inversion of gradient

Example of the recharge identification: surface-GW exchanges

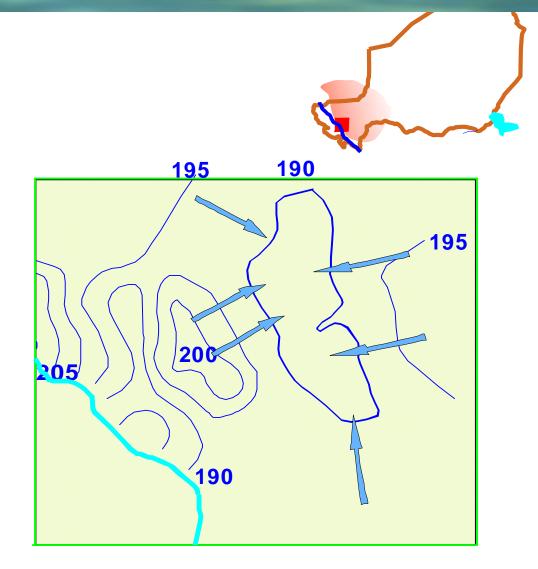






Example of the Continental Terminal aquifer (SW Niger)

focused study over 100*100 km² further validation over a wider area small hydraulic gradients singularity of closed depressions visible annual infitration



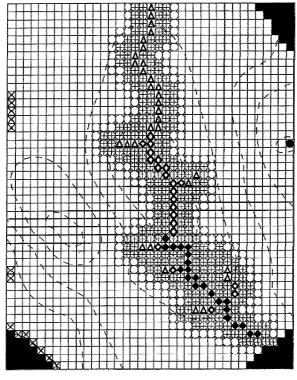
two models:

calibration on a steady state in the 1960s transient simulation from 1960s to 1990s only a small part of the aquifer strong constraint because of the closed depression

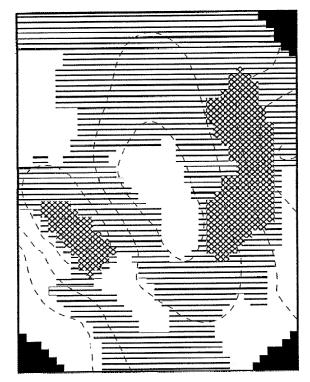
GW model only (G. Favreau, 2000), surface-GW (S. Massuel, 2005)

in both cases, satisfying comparison between measurements and modelled levels

First model (GW only)

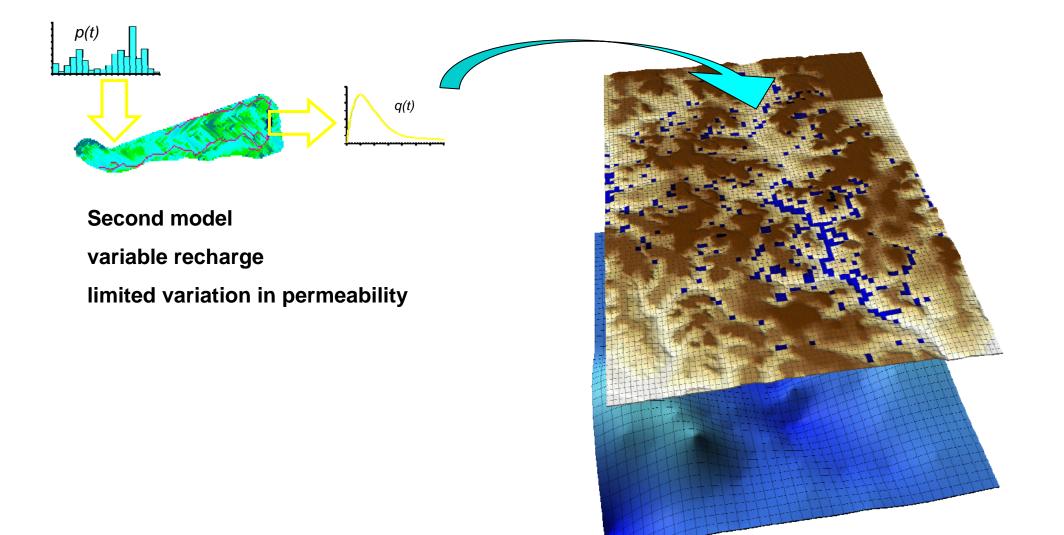


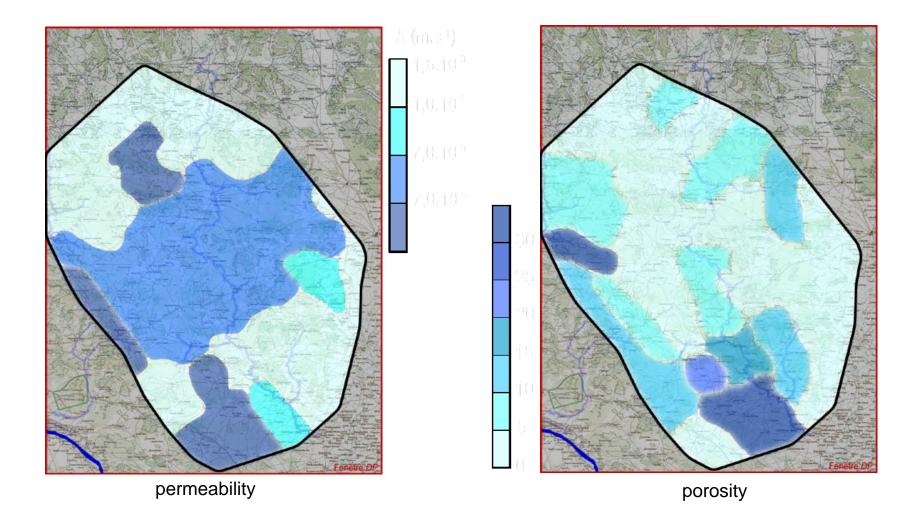
uniform recharge (1 mm/year)



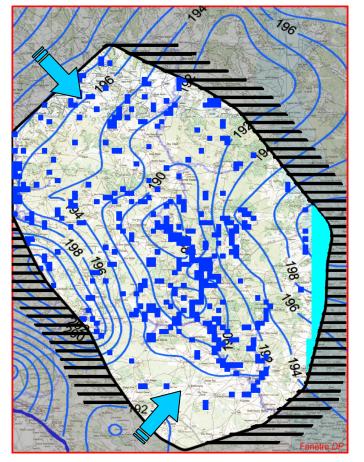
permeability (10⁻⁶ 10⁻³ m/s)

Example of the Continental Terminal aquifer

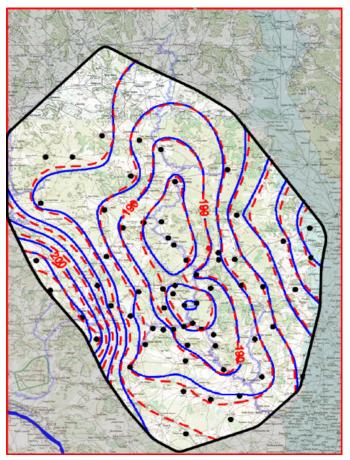




Example of the Continental Terminal aquifer



location of GW recharge areas

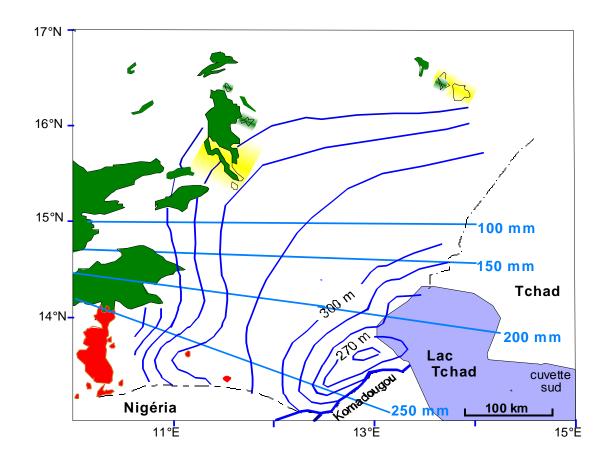


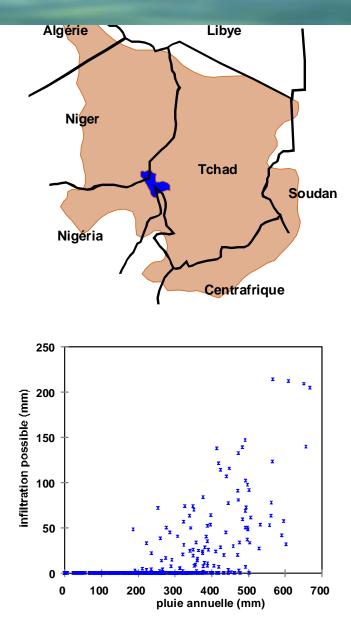
GW levels observed-modelled

In this region anthropization is progressive but fundamental

2 models, with different assumptions and objectives similar quality of modelled GW levels assumptions significantly different

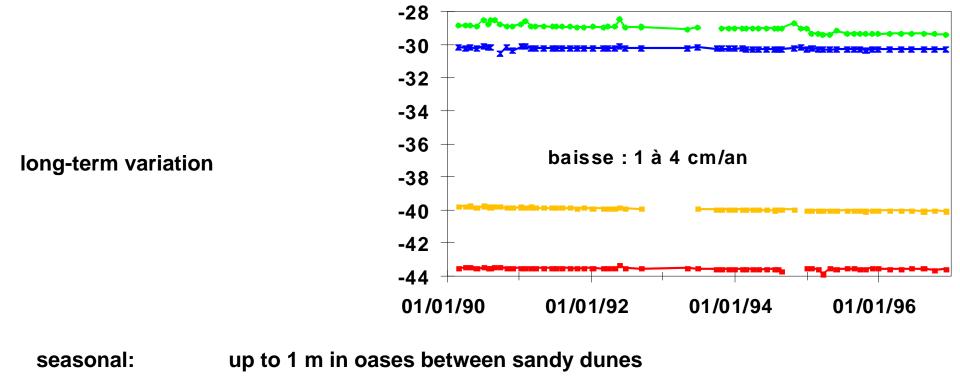
Example of the Lake Chad Basin Quaternay aquifer





infiltration calculated with a tank model with a daily time step

Example of the Lake Chad Basin Quaternay aquifer



variation close to Lake Chad and rivers

no clear variation elsewhere

interannual : variation in inflow from rivers and Lake poor variations elsewhere

heterogeneous and scarce data, unevenly distributed in space and over time GW greatlu constrained by several closed piezometric depressions

various GW models at different scales

- the whole aquifer
- closed depressions
- local sub-units

Example of the Lake Chad Basin Quaternay aquifer example of a general model (M. Leblanc, 2003) XX 100 Kilcenetere Xul XXXXXXX Values of recharge and discharge rates

-0 mm/yr -1 mm/yr -11,5 mm/yr Riech ar ekt => 11 mm/y

U.5 mm/yr 1 mm/yr 2 mm/yr

3 mm/vr

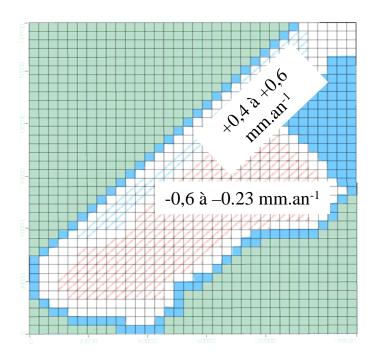
Surface waters

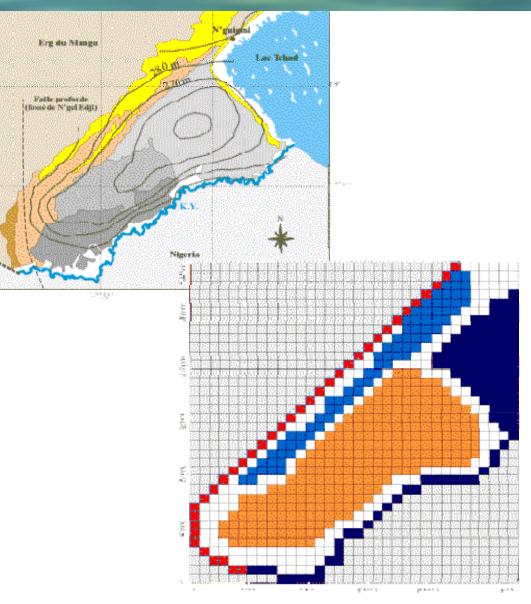
Simulated plazometry

Example of the Lake Chad Basin Quaternay aquifer

example of a sub-regional model:

the Kadzell closed depression (Gaultier, 2004)





Example of the Lake Chad Basin Quaternay aquifer

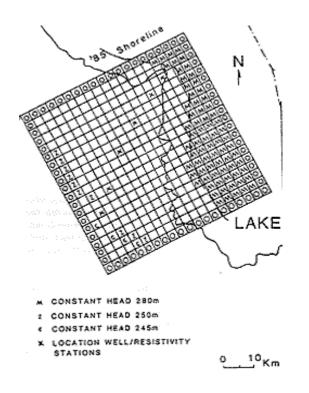
poorly reliable example:

local model in NE Nigeria

unrealistic conditions

heavy assumptions

very limited interest



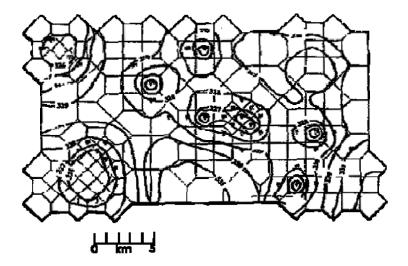
Another example of a local model (NE Nigeria) by Carter

local NE Nigeria : Carter + 50 mm/year

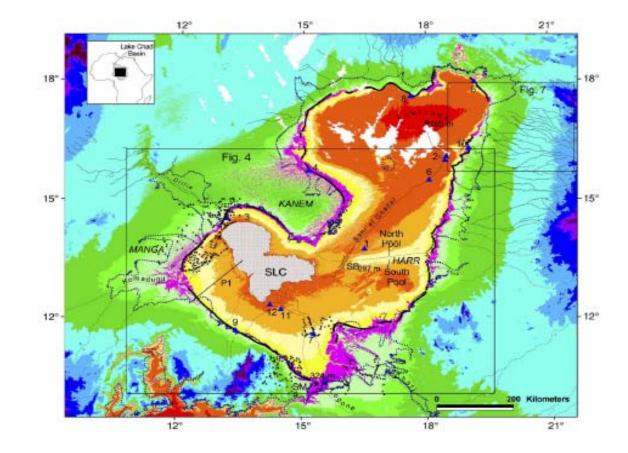
regional Manga Niger : LL + 1 mm/year Gaultier + 0,5 mm/year

whole aquifer : Eberschweiler + 0 Leblanc +/- 1 mm/year

compensation of K and I/O

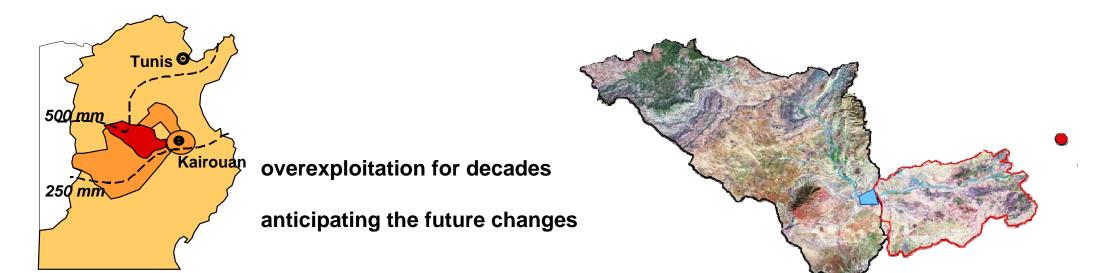


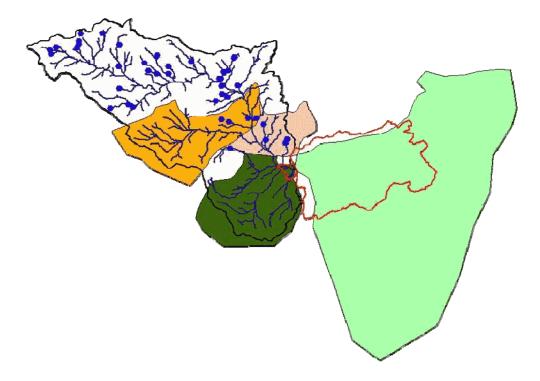
no exchange with the rest of the aquifer recharge 50 mm/year, Evap=5 mm/day K= 2.10^{-4} m/s and S = 38 % considering millenia paléohydrologique

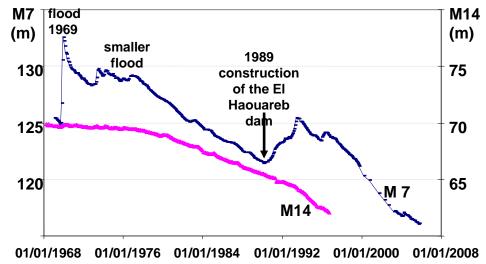


Mega-Chad 6000 years ago

Example of the Merguellil catchment (Kairouan region, central Tunisia)

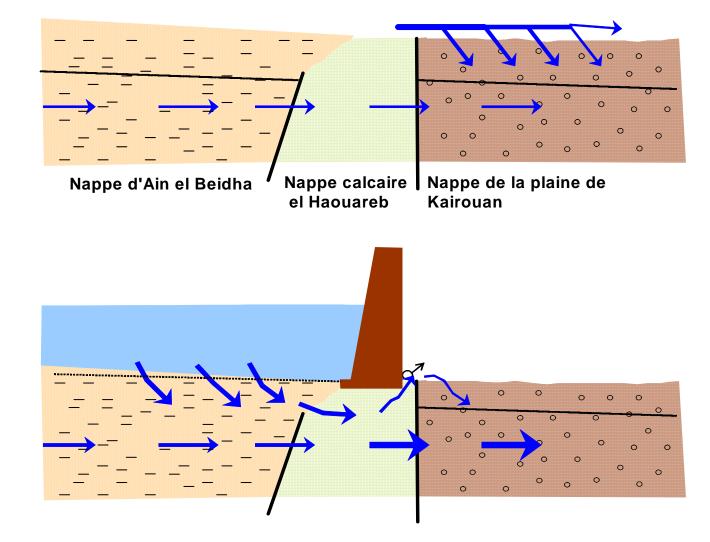






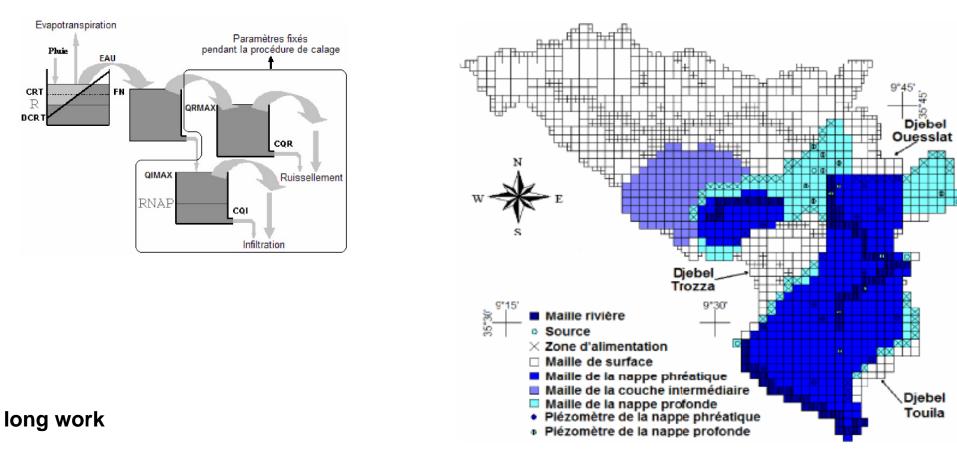
Example of the Merguellil catchment

Before the dam



After the dam

coupled model surface-GW in the upstream catchment (A. Kingumbi)



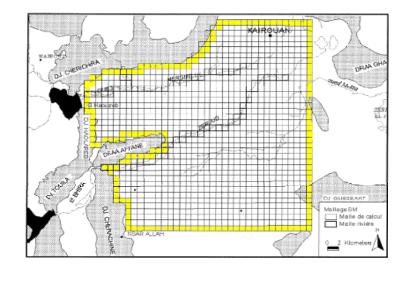
large uncertainty because of environment and data

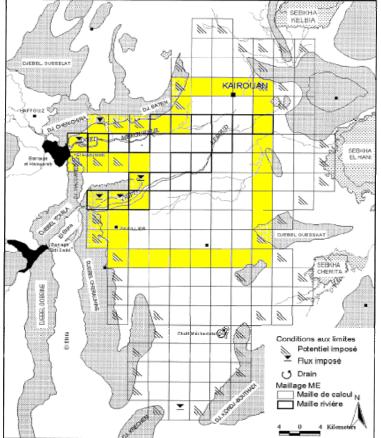
conclusion in contradiction with a parallel work on river flow

Kairouan plain aquifer models keeping the same fundamental assumptions relying on official figures

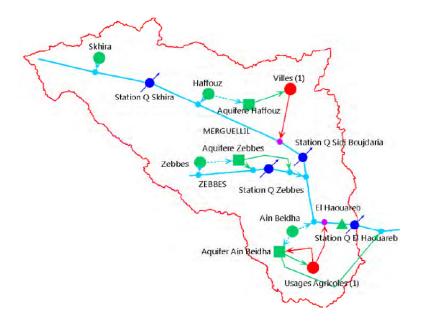
results

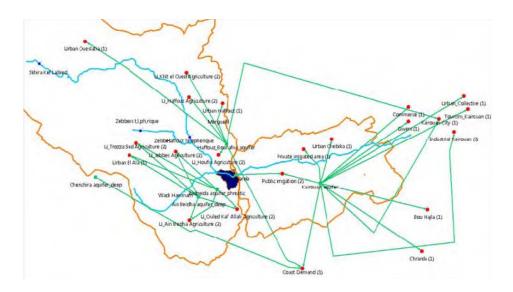
questionable estimate of flows, and then K





GW model later imported in a WEAP model

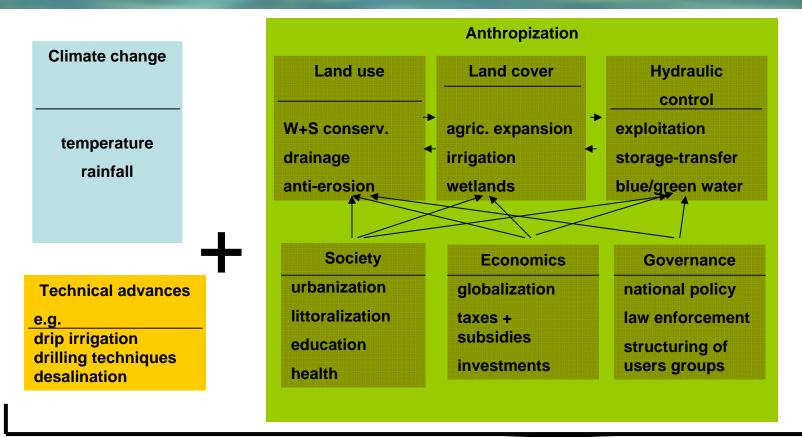


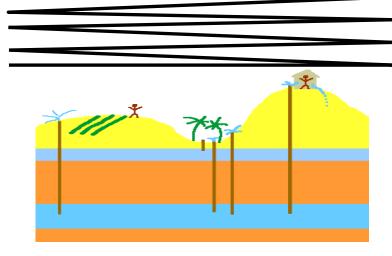


up to now, no consideration of GW quality pollution by agriculture



Drivers of changes in semi-arid aquifers





A major concern is the invisibility of GW

- interpolation by hydrogeologists between sites with information
- different perceptions of GW (in storage and flow) among scientists water users water managers
- no immediate link between an individual action and a change in the GW system
- favouring illegal behaviours (pumping, rejection of contamination)
- political indifference of authorities (no inauguration, no short-term reaction)